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Periodic Water Waves through an Aquatic Forest

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Science

The Asian Tsunami: A Protective Role for Coastal Vegetation

Vaithilingam Selvam, Faizal Parish, Neil D. Burgess, Tetsuva Hiraishi, Vagarappa M. Karunagaran, 3 Michael S. Rasmussen, Lars B. Hansen, Alfredo Quarto, 8 Nyoman Survadiputra9

The scale of the 26 December 2004 Indian Ocean tsunami was almost unprecedented. In areas with the maximum tsunami intensity, little could have prevented catastrophic coastal destruction. Further away, however, areas with coastal tree vegetation were markedly less damaged than areas without. Mangrove forests are the most important coastal tree vegetation in the area and are one of the world's most threatened tropical ecosystems (1).

ling of fluid dynamics suggest that tree vegetation may shield coastlines from tsunami damage by reducing wave amplitude and energy (2). Analytical models show that 30 trees per 100 m2 in a 100-m wide belt may reduce the maximum tsunami flow pressure by more than 90% (3). Empirical and fieldbased evidence is limited, however. Cuddalore District in Tamil Nadu, India,

Measurement of wave forces and model-

provides a unique experimental setting to test

in reducing coastal destruction by tsunamis (4). Cuddalore has a relatively straight shoreline, a fairly uniform beach profile, and a homogenous continental slope. Moreover, the shoreline comprises vegetated as well as nonvegetated areas and was documented by cloudfree pre- and post-tsunami

The force of the tsunami impact in Cuddalore is illustrated by the central part of our study area (Fig. 1). At the river mouth, the tsunami completely destroyed parts of a village (fig. S1) and removed a sand spit that formerly blocked the river. However, areas with mangroves (Fig. 1, dark green polygon) and tree shelterbelts were significantly less damaged than other areas (supporting online text). Damage to villages also varied markedly. In the north, stands of mangroves had five associated villages, two on the coast and three behind the mangrove. The villages on the coast were completely destroyed, whereas those behind the mangrove suffered no destruction even though the waves damaged areas unshielded by vegetation north and south of these villages. In the south, the shore is lined with Casuarina plantations (Fig. 1). Five villages are located within these plantations and all experienced only partial damage. The plantations were undamaged except for rows of 5 to 10 trees nearest to the shore, which were uprooted

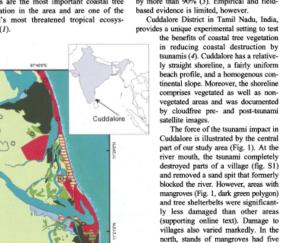


Fig. 1. Pre-tsunami tree vegetation cover and posttsunami damages in Cuddalore District, Tamil Nadu, India.

Open tree vegetation

Dense tree vegetation

Other land cover

Partially damaged

Inundated

Damaged

Water

O Village

BREVIA

Our results suggest that mangroves and Casuarina plantations attenuated tsunamiinduced waves and protected shorelines against damage. Human activities reduced the area of mangroves by 26% in the five countries most affected by the tsunami, from 5.7 to 4.2 million ha, between 1980 and 2000 (5). Conserving

or replanting coas should buffer con events. Mangrov and forestry prod found in artificial Coastal tree vege investments of U (7). Mangroves, h ing only on co which cover ~25 of the Bay of Be servation of dune other tree specie fulfil the same pr

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- P. J. Mumby et a 7. F. Parish, Assessn in Tsunami-Impa Centre, Selangor 8. V. J. Chapman, I of Ecosystems of
- 1977), p. 3. 9. We thank T. Y. C Topp-jørgensen, ance. Supported

Supporting Online I www.sciencemag.or Materials and Metho Figs. S1 and S2

References and Note 10 January 2005; accepted 20 september 2005; 10.1126/science.1118387

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Motivation

2006 West Java Tsunami





Source: Coastal protection in the aftermath of the Indian Ocean tsunami, UN report (2007).

2011 Japan Tohoku Tsunami

Natori, Miyagi Prefecture

(Photographer: Koichiro)





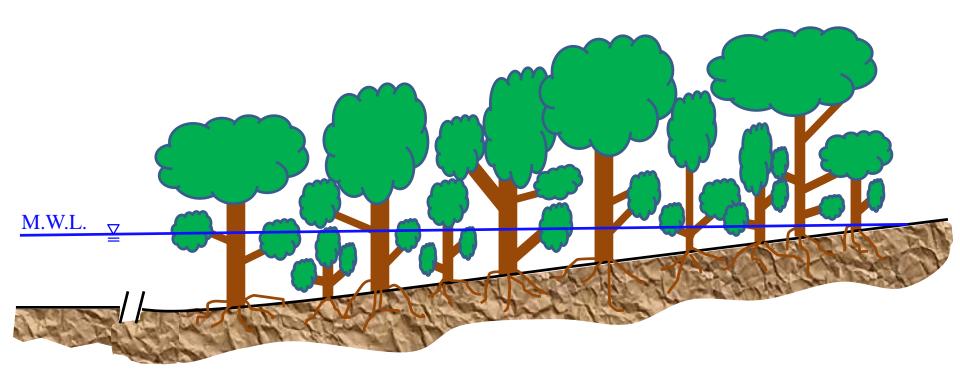
Motivation and Objective

■ Motivation: *Green-belt protection*

Coastal forests as natural barriers against ocean waves

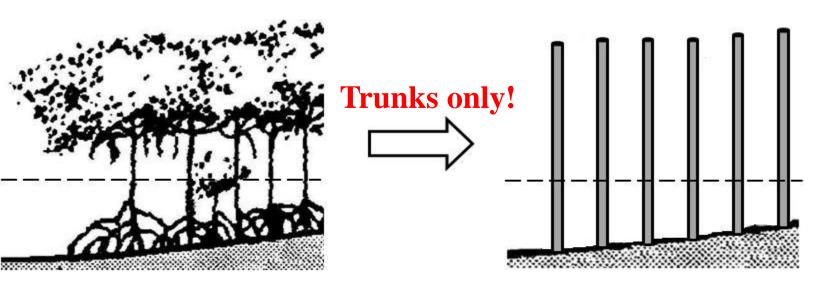
- Wave reflection and wave attenuation by the coastal trees
- Tsunamis, storm surges, tides, daily coastal waves: limited protection!
 - > Focus on dense coastal forests: trees that will stand!
- Coastal ecosystem: dispersion coefficient, etc.
- Objective: To develop a theoretical model for surface waves through emergent trees
 - Emphasize on the propagation and dissipation processes
 - Evaluate the effectiveness of the coastal forests

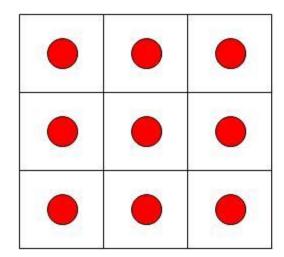
How to describe these coastal trees?

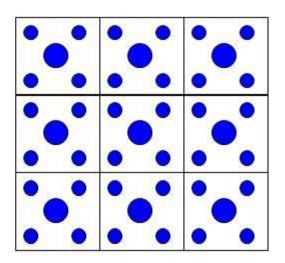


- Distribution of trees?
- Tree trunks? roots? leafs? branches?
- Soil types?
- Trunks: rigid or deformable?
- Emergent or submerged?

An idealized model: rigid cylinders







Horizontal periodicity (top view): cell concept

Mathematical models & simplifications

Basic assumptions

- Slowly varying topography
- Intermediate depth: $kh = \mathcal{O}(1) \Rightarrow 3D$ problem
- Reynolds averaged equations
- Linearized equations: small-amplitude waves
- Eddy viscosity model for wave-forest turbulence

Additional conditions

- \circ Multiple length scales: d, ℓ, h, L
- \circ $d,\ell \ll L \Rightarrow \epsilon = \ell/L = \mathcal{O}(0.1 \mathrm{\ m})/\mathcal{O}(100 \mathrm{\ m}) \ll 1$
- $\circ \ \ell \sim h \text{ or } \ell \ll h$

Intermediate depth

$$kh = \mathcal{O}(1), \ \ell \ll h$$

Two-scale analysis

Series expansions: $\epsilon = \ell/L \ll 1$

$$\mathcal{F} = \mathcal{F}^{(0)} + \epsilon \mathcal{F}^{(1)} + \epsilon^2 \mathcal{F}^{(2)} + \cdots, \quad \mathcal{F} = u_i, w, p$$

$$\mathcal{F} = \mathcal{F}(x_k, X_k, Z, t), \quad k = 1, 2$$

 X_k, Z : macro scales (wavelength L; depth h)

- Homogenization: two distinct scales
 - Horizontal periodicity over wavelength-scale
 - Macro problem: leading-order equations governing wave motions with effects of rigid cylinders
 - Cell (micro-scale) problem: determine the effective parameters due to the presence of a costal forest

Leading-order solutions

Periodic waves

$$F(\vec{x}, \vec{X}, Z, t) = \Re \left\{ \widetilde{F}(\vec{x}, \vec{X}, Z) e^{-it} \right\}$$

Solution forms

$$\widetilde{u}_{i}^{(0)} = -\widetilde{K}_{ij}(\vec{x})\frac{\partial \widetilde{p}^{(0)}}{\partial X_{j}}, \quad \widetilde{w}^{(0)} = -\widetilde{W}(\vec{x})\frac{\partial \widetilde{p}^{(0)}}{\partial Z}, \quad \widetilde{p}^{(1)} = -\widetilde{A}_{j}(\vec{x})\frac{\partial \widetilde{p}^{(0)}}{\partial X_{j}}$$

Macro equations: wavelength-scale

Governing equations for: $\langle \widetilde{u}_i^{(0)} \rangle, \ \langle \widetilde{w}^{(0)} \rangle, \ \widetilde{p}^{(0)}$

$$\langle f \rangle = \frac{1}{\Omega} \int_{\Omega_f} f dx_1 dx_2$$
: averaging over a unite cell Ω

Micro (cell) equations: tree-spacing scale

Governing equations for: $\widetilde{K}_{ij}(\vec{x}), \ \widetilde{W}(\vec{x})$

Homogenization: macro equations

$$\frac{\partial \langle \widetilde{u}_i^0 \rangle}{\partial X_i} + \frac{\partial \langle \widetilde{w}^0 \rangle}{\partial Z} = 0$$
 Eddy viscosity --- to be discuss

--- to be discussed later

$$-\mathrm{i}\langle \widetilde{u}_{i}^{0} \rangle = -n \frac{\partial \widetilde{p}^{0}}{\partial X_{i}} - \left[\oint_{S_{B}} ds \left(-\widetilde{A}_{k} \delta_{ij} + \sigma \left(\frac{\partial \widetilde{K}_{ik}}{\partial x_{j}} + \frac{\partial \widetilde{K}_{jk}}{\partial x_{i}} \right) \right) n_{j} \right] \frac{\partial \widetilde{p}^{0}}{\partial X_{k}}$$

$$\left| -\mathrm{i} \langle \widetilde{w}^0 \rangle = -n \frac{\partial \widetilde{p}^0}{\partial Z} - \left[\sigma \oint_{S_B} \frac{\partial \widetilde{W}}{\partial x_j} n_j \, ds \right] \frac{\partial \widetilde{p}^0}{\partial Z} \right|$$

$$\circ \langle f \rangle = \frac{1}{\Omega} \int_{\Omega_f} f dx_1 dx_2$$
: cell average; n : porosity

- $\circ (\widetilde{K}_{ij}, \widetilde{A}_j, \widetilde{W})$ to be obtained from the cell problem
- \circ No forests: n=1 and $[\cdots]=0$

 \Rightarrow reduce to the common wave equations

Homogenization: cell problem

$$\widetilde{u}_{i}^{0} = -\widetilde{K}_{ij}(\vec{x})\frac{\partial \widetilde{p}^{0}}{\partial X_{j}}, \quad \widetilde{p}^{1} = -\widetilde{A}_{j}(\vec{x})\frac{\partial \widetilde{p}^{0}}{\partial X_{j}}, \quad \widetilde{w}^{0} = -\widetilde{W}(\vec{x})\frac{\partial \widetilde{p}^{0}}{\partial Z}$$

Horizontal directions

$$\frac{\partial \widetilde{K}_{ij}}{\partial x_j} = 0, \quad \vec{x} \in \Omega, \quad -i\widetilde{K}_{ij} = \delta_{ij} - \frac{\partial \widetilde{A}_j}{\partial x_i} + \sigma \frac{\partial^2 \widetilde{K}_{ij}}{\partial x_k \partial x_k}$$

Vertical direction

To be discussed later

$$-i\widetilde{W} = 1 + \sigma \frac{\partial^2 \widetilde{W}}{\partial x_j \partial x_j}, \quad \vec{x} \in \Omega$$

Boundary conditions

Periodicity;
$$\widetilde{K}_{ij} = \widetilde{W} = 0$$
 on S_B ; $\langle \widetilde{A}_j \rangle = 0$

Solution procedure

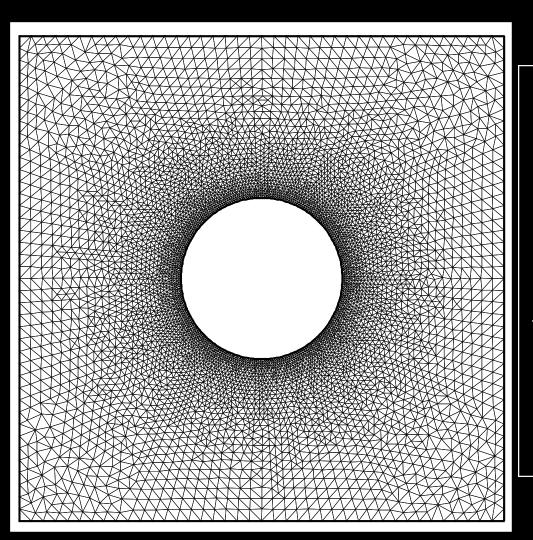
- \circ Solve the cell problem for $\left(\widetilde{K}_{ij},\widetilde{A}_{j},\widetilde{W}\right)$
- Solve the maro equations

$$\frac{\partial}{\partial X_k} \left[\left(n + M_{ik} \frac{\partial \phi}{\partial X_i} \right) \right] + (n+N) \frac{\partial^2 \phi}{\partial Z^2} = 0, \quad i, k = 1, 2$$

$$\begin{cases} \widetilde{p}^{(0)} = i\phi, & \langle \widetilde{u}_{i}^{(0)} \rangle = n \frac{\partial \phi}{\partial X_{i}} + M_{ik} \frac{\partial \phi}{\partial X_{k}}, & \langle \widetilde{w}^{(0)} \rangle = (n+N) \frac{\partial \phi}{\partial Z} \\ M_{ik} = \oint_{S_{B}} ds \left[-\widetilde{A}_{k} \delta_{ij} + \sigma \left(\frac{\partial \widetilde{K}_{ik}}{\partial x_{j}} + \frac{\partial \widetilde{K}_{jk}}{\partial x_{i}} \right) \right] n_{j} \\ N = \sigma \oint_{S_{B}} ds \frac{\partial \widetilde{W}}{\partial x_{j}} n_{j}, & M_{ik}(\vec{X}), N(\vec{X}) = \text{functions of } \nu_{e} \end{cases}$$

Numerical solutions (FDM); Analytical solutions?

Sample cell problem: FEM solutions



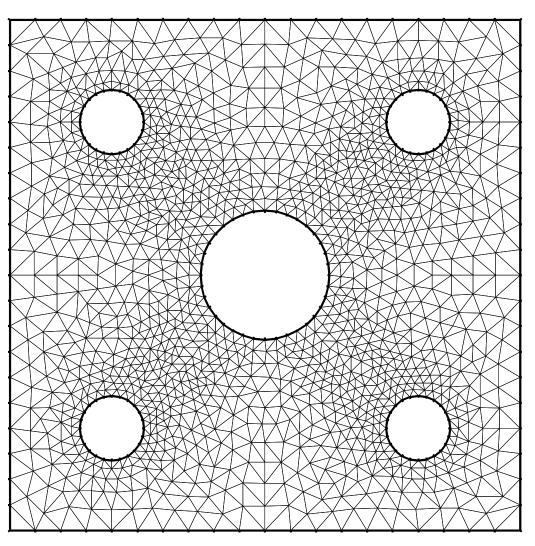
$$\begin{cases} \frac{\partial \widetilde{K}_{ij}}{\partial x_i} = 0 \\ -i\widetilde{K}_{ij} = \delta_{ij} - \frac{\partial \widetilde{A}_j}{\partial x_i} + \sigma \frac{\partial^2 \widetilde{K}_{ij}}{\partial x_k \partial x_k} \end{cases}$$
$$-i\widetilde{W} = 1 + \sigma \frac{\partial^2 \widetilde{W}}{\partial x_j \partial x_j}$$

with

- o Periodicity
- \circ No-slip on the cylinder S_B

$$\circ \langle \widetilde{A}_i \rangle = 0$$

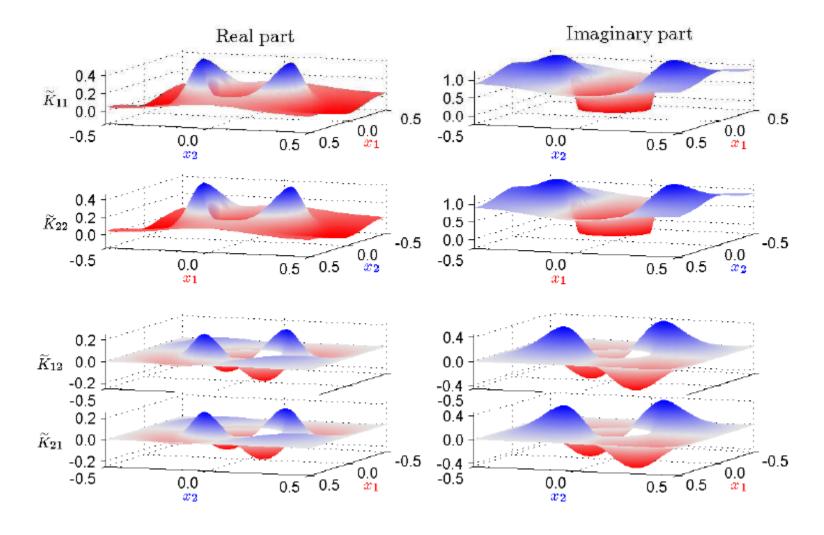
Several cylinders in a unit cell





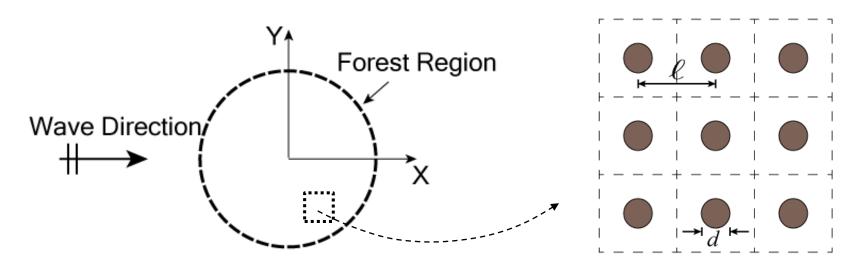


Sample cell solutions



Macro-scale problem for constant eddy viscosity

Period wave through a circular forest



Sketch of the research problem

Arrangement of cylinder array

Leading-order governing equations

(dimensionless)

Forest region:

Complex potential $\tilde{p}^{(0)} = i\phi(X_i, Z)$

$$\frac{1}{r}\frac{\partial}{\partial r}\left(r\frac{\partial\phi^{(F)}}{\partial r}\right) + \frac{1}{r^2}\frac{\partial^2\phi^{(F)}}{\partial\theta^2} + \left(\frac{n+N}{n+M}\right)\frac{\partial^2\phi^{(F)}}{\partial z^2} = 0, \ r < R$$

• Outside region:

$$\frac{1}{r}\frac{\partial}{\partial r}\left(r\frac{\partial\phi^{(I)}}{\partial r}\right) + \frac{1}{r^2}\frac{\partial^2\phi^{(I)}}{\partial\theta^2} + \frac{\partial^2\phi^{(I)}}{\partial z^2} = 0, \ r > R$$

Complex coefficients:

$$M_{ik} = \frac{1}{\Omega} \oint_{S_B} ds \left[-\tilde{A}_k \delta_{ij} + \sigma \left(\frac{\partial \tilde{K}_{ik}}{\partial x_j} + \frac{\partial \tilde{K}_{jk}}{\partial x_i} \right) \right] n_j^{S_B}$$

$$M_{11} = M_{22} \equiv M$$
 ; $M_{12} = M_{21} = 0$

Constant for each cell

$$N = \frac{1}{\Omega} \sigma \oint_{S_B} \left(\frac{\partial \tilde{W}}{\partial x_j} n_j^{S_B} \right) ds$$

Macro-scale analytical solutions

• Forest region:

$$\phi^{(F)} = A_0 \sum_{m=0}^{\infty} \epsilon_m i^m \cos m\theta \left[\sum_{q=0}^{\infty} B_{mq} J_m(\gamma \hat{k}_q r) \frac{\cosh \hat{k}_q(Z+h)}{\cosh \hat{k}_q h} \right]$$

$$\gamma = \sqrt{(n+N)/(n+M)} \qquad \epsilon_m : \text{ Jacobi symbol}$$

• Outside region:

Incident waves

$$\phi^{(I)} = A_0 \sum_{m=0}^{\infty} \epsilon_m i^m \cos m\theta \left\{ \left[J_m(k_0 r) + C_{m0} H_m^{(1)}(k_0 r) \right] \frac{\cosh k_0 (Z+h)}{\cosh k_0 h} \right\}$$

$$+ A_0 \sum_{m=0}^{\infty} \epsilon_m i^m \cos m\theta \left[\sum_{p=1}^{\infty} C_{mp} K_m(\kappa_p r) \frac{\cos \kappa_p (Z+h)}{\cos \kappa_p h} \right]$$

• Matching conditions along r = R

$$\phi^{(I)} = \phi^{(F)}, \quad -h < Z < 0$$

$$\frac{\partial \phi^{(I)}}{\partial r} = (n+M)\frac{\partial \phi^{(F)}}{\partial r}, \quad -h < Z < 0$$

Eddy viscosity by energy balance

- Energy balance concept
 - **➤** Viscous dissipation = Work input
- Wave force on one or more cylinders
 - > A single cylinder: Morison equation

Force =
$$\frac{\rho}{2}C_D u|u|d + \rho C_M \frac{\partial u}{\partial t} \frac{\pi d^2}{4}$$

- ➤ Multiple cylinders (idealized coastal forests)
 - No information on drag and inertia coefficients
 - Focus on time-averaged energy (work-dissipation) balance:

$$\frac{1}{T} \int_{t}^{t+T} \mathcal{E}dt = \overline{\mathcal{E}} = \frac{1}{T} \int_{t}^{t+T} \Psi dt = \overline{\Psi}$$

Time-invariant eddy viscosity

Averaged rate of work done and dissipation

$$\overline{\mathcal{E}} = \frac{1}{2} \rho dC_D \int_{-h}^{0} \overline{\langle u_1^{(0)} \rangle^2 \left| \langle u_1^{(0)} \rangle \right|} dZ + \mathcal{O}(\epsilon)$$

$$\overline{\Psi} = \rho \nu_e \frac{1}{\Omega} \int_{-h}^{0} dZ \iint_{\Omega_f} \frac{1}{2} \left(\frac{\partial u_i^{(0)}}{\partial x_j} + \frac{\partial u_j^{(0)}}{\partial x_i} \right)^2 + \left(\frac{\partial w^{(0)}}{\partial x_i} \right)^2 d\Omega$$

Energy balance

$$\overline{\mathcal{E}} = \overline{\Psi} \implies \frac{\nu_{e}(X)}{\omega a_{0}d} = \frac{\frac{2}{3\pi}C_{D} \left| \langle \widetilde{K}_{11} \rangle \right|^{3} \int_{-h}^{0} \left| \frac{\partial \widetilde{p}^{(0)}}{\partial X} \right|^{3} dZ}{\mathcal{F}_{K} \int_{-h}^{0} \left| \frac{\partial \widetilde{p}^{(0)}}{\partial X} \right|^{2} dZ + \mathcal{F}_{W} \int_{-h}^{0} \left| \frac{\partial \widetilde{p}^{(0)}}{\partial Z} \right|^{2} dZ},$$

$$\mathcal{F}_K = \mathcal{F}_K(\widetilde{K}_{ij}), \quad \mathcal{F}_W = \mathcal{F}_K(\widetilde{W})$$

Bulk (constant) eddy viscosity and drag coefficient

Averaged rate of work done and dissipation

$$\overline{\Psi} = \rho \nu_e \frac{1}{\Omega} \iint_{\text{Forest}} dA_{\text{F}} \int_{-h}^{0} dZ \iint_{\Omega_f} \frac{1}{2} \left(\frac{\partial u_i^{(0)}}{\partial x_j} + \frac{\partial u_j^{(0)}}{\partial x_i} \right)^2 + \left(\frac{\partial w^{(0)}}{\partial x_i} \right)^2 d\Omega$$

$$\overline{\mathcal{E}} = \frac{1}{2} \rho dC_D \iint_{\text{Forest}} dA_{\text{F}} \int_{-h}^{0} \overline{\langle u_1^{(0)} \rangle^2 \left| \langle u_1^{(0)} \rangle \right|} dZ + \mathcal{O}(\epsilon)$$

Energy balance

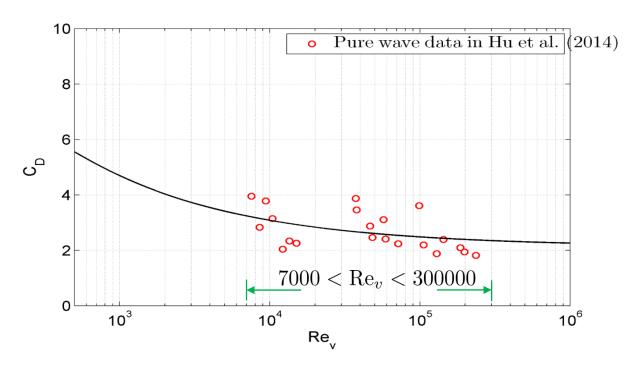
Bulk eddy viscosity
$$\frac{\langle \nu_{e} \rangle}{\langle \omega_{a_0} d} = \frac{\frac{2}{3\pi} \langle C_D \rangle \iint_{\text{Forest}} dA_F \left| \langle \widetilde{K}_{11} \rangle \right|^3 \int_{-h}^{0} \left| \frac{\partial \widetilde{p}^{(0)}}{\partial X} \right|^3 dZ}{\iint_{\text{Forest}} \left[\mathcal{F}_K \int_{-h}^{0} \left| \frac{\partial \widetilde{p}^{(0)}}{\partial X} \right|^2 dZ + \mathcal{F}_W \int_{-h}^{0} \left| \frac{\partial \widetilde{p}^{(0)}}{\partial Z} \right|^2 dZ \right] dA_F}$$

Hu et al. (2014) for bulk drag coefficient:

- Conduct flume tests on a forest belt with finite thickness 6 m (normal incidence)
- ➤ Use rigid cylinders to model the forest
- ➤ Directly measure the forces on individual cylinders
- ➤ Use the same time-averaged formula to calculate the drag coefficient
- > Use both submerged and emergent cylinders
- ➤ Propose a new drag coefficient formula for periodic waves and combined wave-current conditions

Bulk drag coefficient

■ Hu et al. experiments (2014)



$$\langle C_D \rangle = \frac{50}{\text{Re}_v^{0.43}} + 2.13 \left[1 - \exp\left(-\frac{\text{Re}_v}{120.74}\right) \right]$$

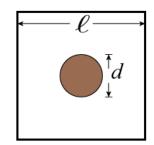
> Reynolds number

$$Re_v = \frac{U_{\text{mid}}r_v}{\nu}$$

> Hydraulic radius

$$r_v = \frac{\pi}{4}d\left(\frac{n}{1-n}\right)$$

> Porosity



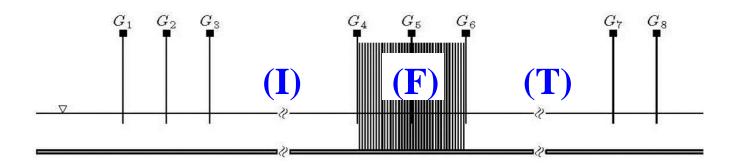
$$n = 1 - \frac{\pi d^2/4}{\ell^2}$$

Iterative scheme for eddy viscosity

- Eddy viscosity is considered as a constant in the entire forest.
- Iteration for eddy viscosity
 - An initial guess for eddy viscosity
 - Solve the cell problem first
 - Solve the macro problem based on the cell solutions
 - Calculate the new eddy viscosity
 - Convergence test: around 10 iterations

Model testing & application

Intermediate depth: a 2HD example



$$\phi^{(I)} = A_0 \frac{\cosh k_0 (Z+h)}{\cosh k_0 h} e^{ik_0 X} + A_0 R \frac{\cosh k_0 (Z+h)}{\cosh k_0 h} e^{-ik_0 X} + \sum_{p=1}^{\infty} A_p \frac{\cos \kappa_p (Z+h)}{\cos \kappa_p h} e^{\kappa_p X}$$

$$\phi^{(F)} = \sum_{q=0}^{\infty} \frac{\cosh \hat{k}_q (Z+h)}{\cosh \hat{k}_q h} \left[C_q e^{i\gamma \hat{k}_q X} + D_q e^{-i\gamma \hat{k}_q X} \right]$$

$$\phi^{(T)} = A_0 T \frac{\cosh k_0 (Z+h)}{\cosh k_0 h} e^{ik_0 X} + \sum_{p=1}^{\infty} B_p \frac{\cos \kappa_p (Z+h)}{\cos \kappa_p h} e^{-\kappa_p X}$$

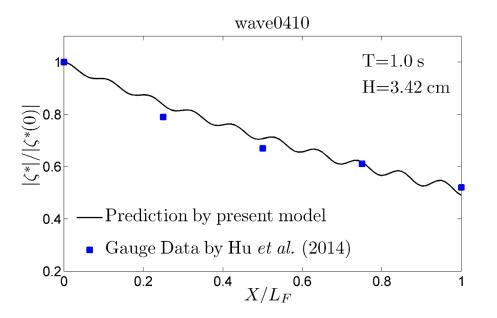
 $(R, T, A_p, B_p, C_q, D_q)$: Obtained by matching velocity & pressure along interfaces

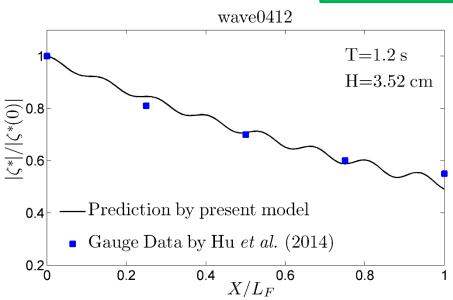
Dispersion relationships:

$$1 = k_0 \tanh k_0 h \; ; \; 1 = -\kappa_p \tan \kappa_p h \; (p \ge 1) \; ; \qquad 1 = \left(\frac{n+N}{n}\right) \hat{k}_q \tanh \hat{k}_q h, \quad (q \ge 0)$$

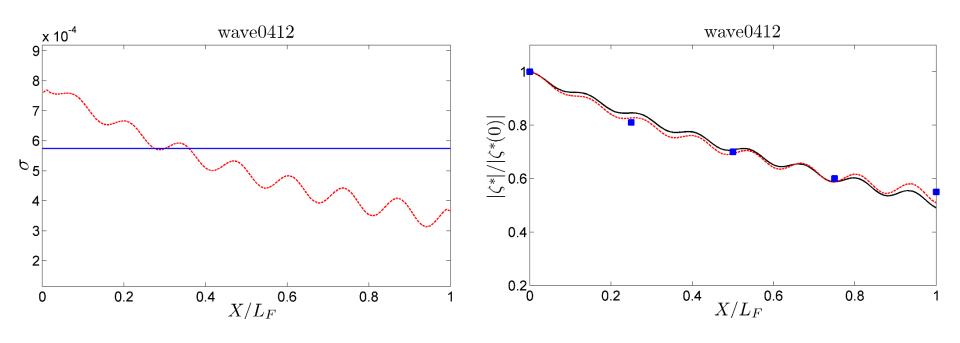
Model validation (Hu et al. 2014)

- Two-dimensional forest belt with normal incidence
 - Compare the dimensionless relative wave amplitude inside the forest with the wave gauge measurements





How appropriate is the constant eddy viscosity model?



Bulk eddy viscosity vs.
Varying eddy viscosity

Normalized dimensionless wave amplitude

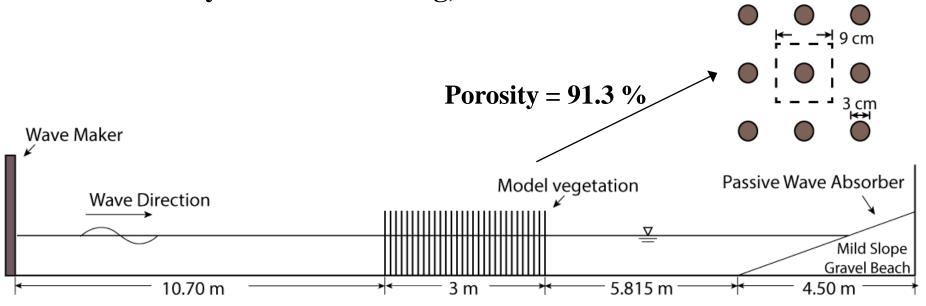
NEW LABORATORY EXPERIMENTS

UNIVERSITY OF CANTABRIA, SPAIN

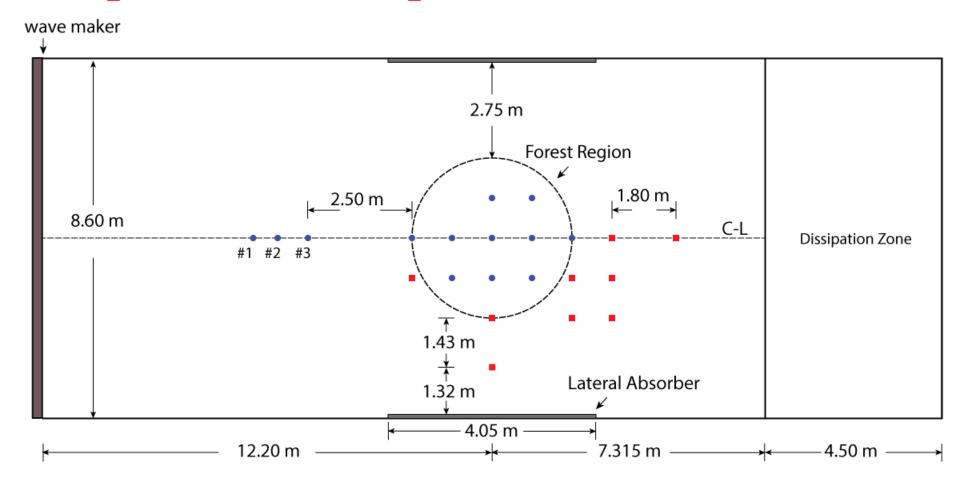
Experimental setup

- Wave basin: 28 m long, 8.6 m wide and 1.2 m deep
- Wave maker: piston type with 10 independent paddles
- Passive wave absorber: mild slope gravel beach (1:12)
- Forest region: circle with 3 m in diameter, 880 cylinders in use

• Circular cylinder: 50 cm long, 3 cm in diameter

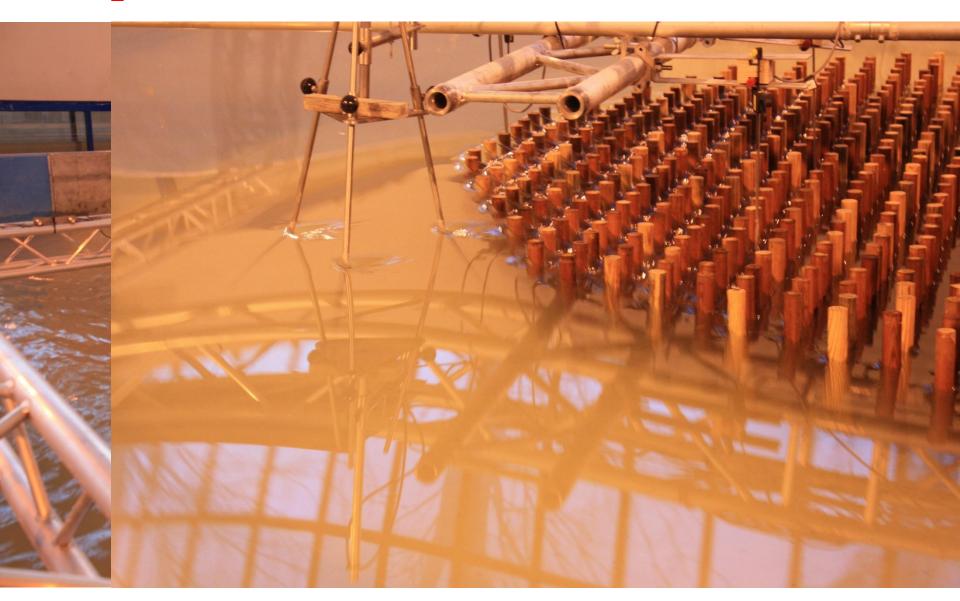


Experimental setup (cont.)



• Instrumentation: 22 wave gauges

Experiments



Experimental conditions

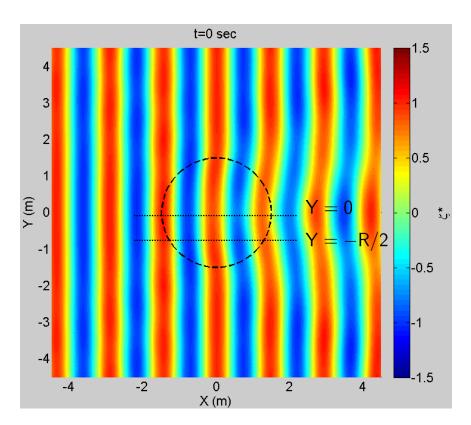
- Two water depths were tested (h=30 cm, 40 cm)
 - Wave period ranges from 1.00 sec to 2.75 sec
 - Incident wave height ranges from 2.50 cm to 7.56 cm

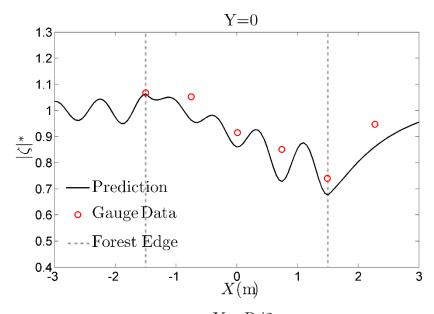
Table 1. Frequency Test II

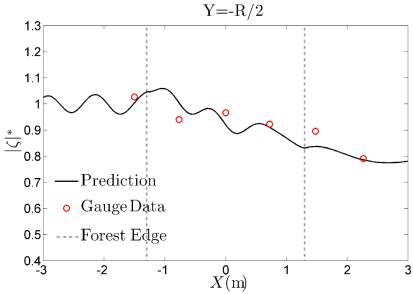
Case	h (cm)	T(s)	H (cm)	L (m)	kh	kA	Re _v	C _D	**
4F2	40	1.00	4.90	1.464	1.717	0.105	9829	3.088	7.04E-04
4F3		1.25	5.54	2.052	1.225	0.085	13361	2.970	1.28E-03
4F4		1.50	5.38	2.616	0.961	0.065	14275	2.946	1.85E-03
4F5		1.75	5.30	3.162	0.795	0.053	14768	2.934	2.50E-03
4F6		2.00	5.22	3.695	0.680	0.044	15017	2.928	3.21E-03
4F7		2.25	5.04	4.220	0.596	0.038	14902	2.931	3.88E-03
4F8		2.50	5.06	4.739	0.530	0.034	15192	2.924	4.73E-03
4F9		2.75	5.22	5.254	0.478	0.031	15744	2.912	5.78E-03

Numerical results vs. data (a)

	Water depth	0.40 m		
4E2	Wave period	1.00 s		
4F2	Wave length	1.46 m		
	Wave height	4.90 cm		

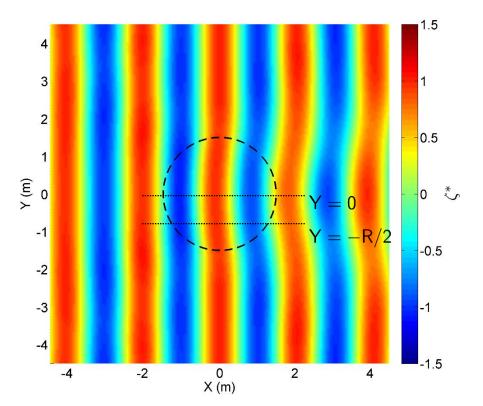


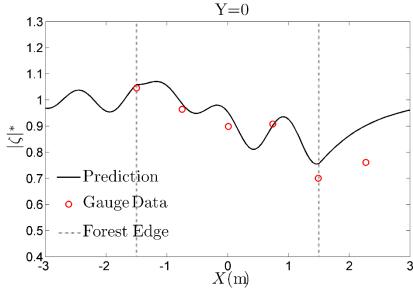


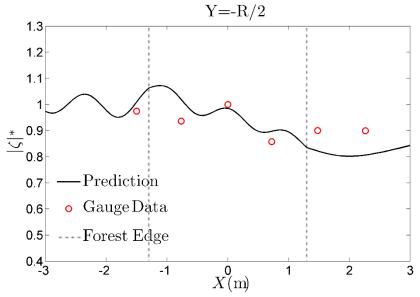


Numerical results vs. data (b)

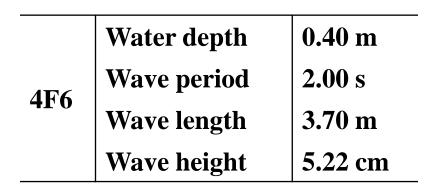
	Water depth	0.40 m		
4E2	Wave period	1.25 s		
4F3	Wave length	2.05 m		
	Wave height	5.54 cm		

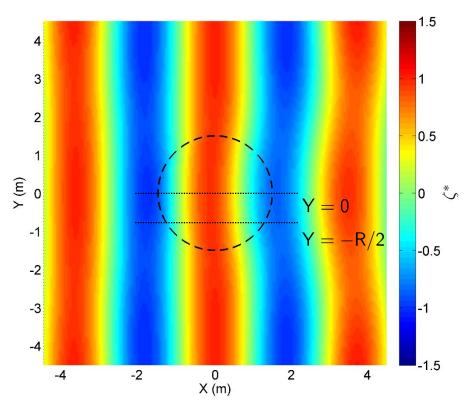


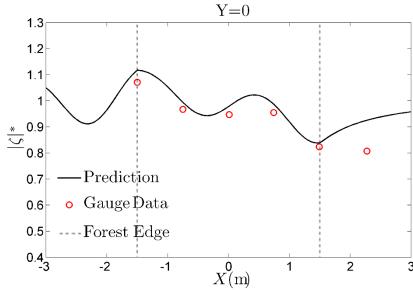


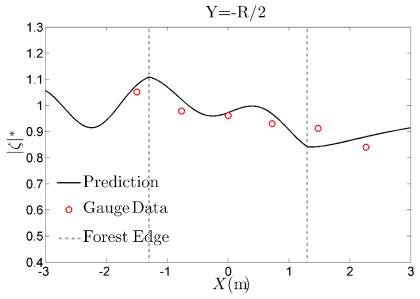


Numerical results vs. data (c)









Surface waves through emergent coastal trees

Summary

- A macro-theory for wave motions, with effective coefficients obtained numerically from the micro-scale problem
- Analytical and numerical solutions have been discussed
- New eddy viscosity model
- o Good agreements between the theory and the experimental data

■ Important facts

- Strong wave attenuation
- Considerable reflected waves
- Theory can be used as a design guideline

■ Improvements

- Choice of eddy viscosity: need more experimental works!
- Weakly wave nonlinearity

End.

Thanks!

Questions?